Understanding the anvil cloud area feedback



Brett McKim, Sandrine Bony, Marion Saint-Lu, Jean-Louis Dufresne, Claudia Stubenrauch LMD | EUREC4A Meeting | 8 December 2022

University of Exeter

EULBRIGHT

Image: https://eol.jsc.nasa.gov/SearchPhotos/photo.pl?mission=ISS066&roll=E&frame=37532

Knowns and unknowns in climate feedbacks



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• Climate models simulate the longwave clear-sky feedback to within 5% $\lambda_{cs} = 1.88 \pm 0.09 \text{ Wm}^{-2} \text{K}^{-1}$

- Cloud feedback uncertainty is larger than the feedback itself $\lambda_{clouds} = 0.45 \pm 33 \text{ Wm}^{-2} \text{K}^{-1}$
- Much uncertainty comes from the anvil cloud area feedback $\lambda_{iris} = -0.2 \pm 0.2 \text{ Wm}^{-2} \text{K}^{-1}$





Zhang et al, 2020; Sherwood et al, 2020

Uncertainties in the anvil cloud area feedback

2 distinct questions

- How much does anvil cloud fraction change with warming?
- What is the the radiative feedback due to that change?





• Models of RCE simulate a diversity of f_h changes—some positive, some negative.

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- This diversity might stem from microphysical parameterizations. $f_h = \text{CSC} \cdot \text{microphysical tendencies}$
- It is a problem that models can't agree on this aspect of climate change.



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• This change ultimately comes from decreased static stability with warming (outweighing the increased radiative cooling)

$$f_h \propto \text{CSC} = \partial_z w_{\text{sub}} = \partial_z \left(\frac{\mathcal{H}}{\Gamma_d - \Gamma} \right)$$

"Stability Iris" (Bony et al, 2016)



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• Do anvil area changes counteract changes in surface temperature?



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Credit Rene Magritte, "The False Mirror," 1929 O 2020 C. Herscovici / Artists Rights Society (ARS), New York

What is the the radiative feedback due to a decreasing f_h?

• The anvil cloud area can be estimated with the change in cloud radiative effect (CRE) with warming, but the interpretation is complicated.



Boundary layer buoyancy, B (10⁻², nondimensional)



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- Clouds and clear-skies change
- Other cloud properties change
- Different reference response than traditional feedback decompositions



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- Clouds and clear-skies change
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- Different reference response than traditional feedback decompositions
- Can we approach this problem in another, simpler way?



Boundary layer buoyancy, B (10⁻², nondimensional)







One Grid Cell



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$$\lambda = \frac{dN}{dT_s} = \frac{d[S - R]}{dT_s}$$



+...(low cloud feedbacks)



+...(low cloud feedbacks)

A number of questions arise:

- What can we learn from this equation?
- How do we validate it?
- Can we estimate the iris feedback?
- What is missing from or assumed in this model?

$$\lambda_{\rm iris} = \frac{1}{f_h} \frac{df_h}{dT_s} \Big[{\rm CRE}_h + \lambda_{cs} (T_s - T_\ell) f_\ell f_h + S^{\downarrow} (1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h \Big]$$

The iris feedback owes to 3 contributions:

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1. The fractional change in high cloud area



longwave

shortwave

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The iris feedback owes to 3 contributions:

- 1. The fractional change in high cloud area
- 2. The high cloud radiative effect in the absence of low clouds



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The iris feedback owes to 3 contributions:

- 1. The fractional change in high cloud area
- 2. The high cloud radiative effect in the absence of low clouds
- 3. The radiative effect of low clouds *no longer* blocked by the high clouds



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 $CRE_h \approx -4 \text{ Wm}^{-2} \text{K}^{-1}$



Hartmann et al, 2017

$$\lambda_{\rm iris} = \frac{1}{f_h} \frac{df_h}{dT_s} \Big[{\rm CRE}_h + \lambda_{cs} (T_s - T_\ell) f_\ell f_h + S^{\downarrow} (1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h \Big]$$

 $\lambda_{cs}(T_s - T_\ell) f_\ell f_h \approx -2 \ \mathrm{Wm}^{-2} \mathrm{K}^{-1} \cdot 15 \ \mathrm{K} \cdot 0.1 \cdot 0.15 \approx -0.5 \ \mathrm{Wm}^{-2}$ $S^{\downarrow}(1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h \approx 340 \mathrm{Wm}^{-2} \cdot 0.87 \cdot 0.5 \cdot 0.1 \cdot 0.5 \cdot 0.15 \approx 1.1 \ \mathrm{Wm}^{-2} \mathrm{K}^{-1}$





Bony et al, 2015; Bony et al, 2005

$$\langle \lambda_{\rm iris} \rangle = \frac{\langle \overline{A_h} \rangle}{\langle \overline{f_h} \rangle} \frac{d \langle f_h \rangle}{d \langle T_s \rangle} \Big[\langle \overline{\text{CRE}_h + \lambda_{cs} (T_s - T_\ell) f_\ell f_h} + S^{\downarrow} (1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h} \rangle \Big]$$

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• What do the terms mean?



- Tropical average (+/- 30 deg N)

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- What do the terms mean?
 - $\langle \cdot \rangle$ T
 - Tropical average (+/- 30 deg N)
 - Temporal average over all years
- Let's validate using ...
 - HadCRUT for T_s, ERA5 for atmospheric T
 - CALIPSO for f_h and f_l , CERES for CRE and surface albedo
 - Estimate clear-sky feedback as -2
 - Estimate cloud albedo by fitting predictions of *S* and *R*







Following Saint-Lu et al, 2020: Use cloud fraction of clouds with optical depth 0.3 < tau < 5

* This definition of anvils excludes clear-sky regions, as well as subvisible cirrus clouds, and the cores of deep convective clouds.

MONTHLY SNAPSHOT

10 YEAR AVERAGE



- Each circle is a different year (averaged from July to June)
- We get a decrease in f_h over the 10-year record, consistent with Saint-Lu et al, 2020









0.3

0.2

 $R = R_{cs}(1 - f_h)(1 - f_\ell)$ $+ \sigma T_h^4 f_h + R_{cs}|_{T_\ell}(1 - f_h) f_\ell$ $S = S^{\downarrow}(1 - \alpha_h f_h)(1 - \alpha_l f_l)(1 - \alpha_s)$

- We don't have cloud albedo data, so we tune it
- Fit albedo to minimize mean CRE error



"With four parameters I can fit an elephant, and with five I can make him wiggle his trunk"

- John von Neumann

$$\langle \lambda_{\rm iris} \rangle = \frac{\langle \overline{A_h} \rangle}{\langle \overline{f_h} \rangle} \frac{d \langle f_h \rangle}{d \langle T_s \rangle} \Big[\langle \overline{\text{CRE}_h + \lambda_{cs} (T_s - T_\ell) f_\ell f_h} + S^{\downarrow} (1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h} \rangle \Big]$$

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$$0.26 + -0.37 + 1.07$$

Predicted values {'iris_feedback': -0.05 'cre_h': 0.26 'lowcloud_lw': -0.37 'lowcloud_sw': 1.07 'cf_c': '7' 'albedo': 0.4 'fh_avg': 0.17 'df_dts': -0.02 'area_average': 0.50



Summary

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- Our simple model describes how the iris feedback is a competition between a small CRE_h and small low cloud unmasking effect.
- We use observations and the model to predict a nearly neutral iris feedback.



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- Our simple model describes how the iris feedback is a competition between a small CRE_h and small low cloud unmasking effect.
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Questions and Follow ups

- We miss middle clouds, thin cirrus, etc. Does that matter? Does the optical depth considered matter?
- Error/uncertainty analysis?
- Can we use this model to understand intermodel spread in the anvil area feedback?
- Can we understand why Lindzen et al 2001 thought the feedback was so important?
- Why is CRE_h small in the first place?
- Do low clouds ever matter for the iris feedback?
- Can we understand if the stability iris radiative effect influences low cloud cover or the circulation?

$$\begin{split} \mathbf{A}_{\mathrm{iris}} &= \frac{1}{f_h} \frac{df_h}{dT_s} \Big[\mathrm{CRE}_h + \lambda_{cs} (T_s - T_\ell) f_\ell f_h + S^{\downarrow} (1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h \Big] \\ & \quad \text{high cloud} \\ & \quad \text{radiative effect} \\ & \quad \text{CRE}_h^{\mathrm{LW}} \\ & \quad \text{CRE}_h^{\mathrm{LW}} \\ \end{split} \quad \begin{array}{l} \text{low cloud} \\ & \quad \text{low cloud} \\ & \quad \text{adiative effect} \\ & \quad \text{CRE}_h^{\mathrm{SW}} \\ \end{array} \quad \begin{array}{l} \text{low cloud} \\ & \quad \text{unmasking} \\ & \quad \text{CRE}_h^{\mathrm{SW}} \\ \end{array} \quad \begin{array}{l} \text{low cloud} \\ & \quad \text{unmasking} \\ & \quad \text{S}^{\downarrow} (1 - \alpha_s) \alpha_\ell f_\ell \alpha_h f_h \Big] \\ \end{array} \end{split}$$



longwave

shortwave

Other directions

- Apply to low clouds feedbacks by distinguish cloud types, introducing cloud-controlling factors, etc.
- Apply to forcings, i.e. "cloud masking" (*a-la* Jeevanjee et al, 2021)
- Motivate a new feedback decompositions that more clearly distinguish clear sky and cloud feedbacks (*a-la* Yoshimori et al, 2020)
- Use to calculate priors more objectively in ECS estimates (by diagnosing the feedbacks from observations)

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